Rayleigh lidar temperature profiles between 15 - 60 km during OZITOS campaign in Río Gallegos (51° 55'S, 69° 14'W), Argentina

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SUMMARY
The determination of temperature measurements from the Rayleigh scattering is an important remote sensing technique for obtaining stratospheric profiles. This technique is applied to signals acquired by a Rayleigh lidar (Light Detection and Ranging). Currently the Observatorio Atmosférico de la Patagonia Austral (51° 55’S, 69° 14’W) in Río Gallegos, Argentina is part of the UVO3 Patagonia project in collaboration with the laboratory of Ozone and UV Radiation in the city of Punta Arenas, Chile. The objective of this paper is to do an introduction to temperature profiles using a Rayleigh lidar as a part of the DIAL system. In CEILAP group basically can obtain ozone profile using a DIAL system described (Wolfram et al., 2008). The LabO3RUV measured using ECC balloon sonde (Electrochemical Concentration Cell) was designed by Komhyr (Komhyr 1969, 1971). The final objective of this paper is to do an introduction to temperature profiles using a Rayleigh lidar which will be described below. Also a campaign of ozone sounding made in Río Gallegos in March 2011, called OZITOS II (OZone profile aT RIO GallegOS) was carried out in March 2011 in Río Gallegos with the temperature profile retrieved by the Rayleigh lidar using the line of 355 nm, in the same period. The results presented in this paper are validated through intercomparisons with measurements made by MLS instrument (Microwave Limb Sounder) onboard the NASA AURA satellite platform and NCEP data.

Key words: Rayleigh lidar, temperature profile, radiosounding measurements

INTRODUCTION
The lidar emerged as a powerful technique for the remote sensing of the atmosphere. The Rayleigh scattering due to air molecules has been widely used over the past 20 years to determine the temperature profile of the atmosphere between 30 and 90 km altitude. This method allows to study the dynamics of the middle atmosphere with high vertical resolution and temporal evolution. The Rayleigh lidar is limited by aerosol scattering, ozone absorption, and dense atmospheric attenuation. To overcome these difficulties, the wavelength dependent non-elastic Raman scattering technique has been employed recently (Gross et al., 1997) (Gu et al., 1997) (Nedeljkovic et al., 1993). However, Raman lidar requires a high-power laser transmitter to improve the low-level signal conditions because the Raman scattering cross section is about 3 orders of magnitude smaller than that of the Rayleigh scattering.

The determination of temperature profiles using a Rayleigh lidar which will be described below. Also a campaign of ozone sounding made in Río Gallegos in March 2011, called OZITOS II (OZone profile aT RIO GallegOS) was carried out in March 2011 in Río Gallegos with the temperature profile retrieved by the Rayleigh lidar using the line of 355 nm, in the same period. The results presented in this paper are validated through intercomparisons with measurements made by MLS instrument (Microwave Limb Sounder) onboard the NASA AURA satellite platform and NCEP data.

The methodologies described in this section were separated in two parts: the first one describe how obtain a temperature profile from a Rayleigh lidar as a part of the DIAL system. The second one, tried describe the sensor used for the validation of temperature profile from Rayleigh lidar.
Rayleigh lidar temperature profiles

Lidar temperature measurements require that only molecular Rayleigh scattering contributes to the return signal and Mie scattering from aerosols is negligible. This is usually the case above 30 km, even after a volcanic eruption such as Mt. Pinatubo (Steinbrecht and Carswell, 1995). When the Mie scattering is not negligible which occurs typically below 30 km, the temperature value is lower than the real one due to the effects of aerosols.

The temperature algorithm only Rayleigh-scattered light signals produced by the atmosphere from the third harmonic of Nd-YAG laser at 355 nm were used. During the lidar measurements, the output of the multi-channel counters (MCS) provides the raw data as single ASCII files, with an integration time of 1 minute. The retrieval algorithm reads two raw data sets at 355 nm (high and low sensitivity), then performs a data integration variable from 1 to 3 hours. In the next step, two corrections are applied to remove systematic errors in the signals: background signals, Signal-Induced Noise (SIN). The objective of these corrections is to obtain a pure lidar backscattering signal. Then both corrected signals are merged by means of linear fitting in the 20-25 km range. After this corrections, we retrieved the temperature profile from the lidar signal.

ECC sondes

The balloon sondes used during OZITOS II campaign are configured by a radiosounding and an ECC which is the responsible for the detection of ozone concentration. In our experiment the ECC sonde launched has also a radiosounding which can measure temperature, humidity and pressure. The radio receptor is a Lockheed Martin LMG6. It was used for store all data emitted by the sonde. As sensor we used a meteorological radiosounding LMS6. An ECC model EN-SCI Corporation was used for measure the ozone concentration.

**OZITOS II CAMPAIGN**

In December 2008, the instrument DIAL for the measurements of stratospheric ozone profile deployed in the Patagonian city of Río Gallegos was accepted as part NDACC (Network for the Detection Atmospheric Composition Change). This new stage of the instrument must satisfy new requirements as intercomparisons with other kind of sensor to check the stability and guarantee a quality in the measurements. Very often different groups around the world used ECC balloon sondes for measured ozone concentration in a region between the surface up to 30 km aprox. Though the principal objective was to make validations between DIAL and ECC balloon sondes, this paper showed the comparison between temperature profile derived with the 355 nm line as described above and the temperature profile measured with the radiosounding, during OZITOS II campaign.

**Experimental design**

The night March 17, 2011 both groups decided to lunch in a same night three balloon sonde in coincidence with the DIAL operation. The aim was study the minimum time of integration in the data files acquired by DIAL systems. The schedule of the experimental design is showed in Figure 1.

Where the horizontal red bar indicate the total time of measurement of the DIAL system and the blue bar indicate the time of flight of the ozonesounding. Can you show that the DIAL systems was operating more than nine hours. If we select a time of integration of three hours we can obtain three independent measurements that each one can be compared with each balloon sonde launched (same period time of flight).

Now we can decrease the time of integration, and we can derivate for example more profiles. The advantage using signals obtained by a DIAL system, is that we can use the signals in 355 nm from the Nd:YAG laser for retrieved a temperature profile without produced any interference on the ozone measurements.

**Results and discussion**

We have taken from the total measurement about nine hours from Rayleigh, three independent period of time which we calculated the temperature profile using a time of 180 minutes of integration. This time is quasi-coincident with the time of flight of the ozonesounding launched beside, 1 km away of the DIAL system in Río Gallegos. This means that we can compare temperature derived from both instruments.

In Figure 2, we showed the comparison of the temperature profile between the Rayleigh lidar temperature and the radiosounding.
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Figure 2. Comparison of the temperature profiles; a), b), c) are the temperature profile by the Rayleigh lidar (red line) compared with radiosounding (blue line). b) show the comparison with NCEP data (dashed green line with square) for March 17, 2011 in Río Gallegos and c) show the comparison with a MLS sensor (dashed black line with square) aboard AURA-NASA in March 17, 2011 lat=-51.74 °, lon= -75.23 °, time: 06:03:09 (UTC).

The region of comparison between both instruments is a disadvantage due they have different heights of cover. In the case of the Rayleigh lidar we can obtain temperature profile from aprox. 10 km up to 60 km. While in the balloon sonde only we can measure temperature profile between the surface up to 32 km. The effective zone where both instruments can be compared, cover the range 10 to 32 km, aprox, limited in the lower part for the lidar and in the higher part for the balloon burst altitude.

The Figure 2 has shown the good agreement between the different profiles, having a relative error lidar - radiosounding lower than 4 %. Additional in b) we superposed the NCEP data for the same day of measurement, and c) show the comparison with the data provided by the MLS instrument aboard the plataforma AURA-NASA.

CONCLUSION

This paper has shown three independent temperature profiles derived with a Rayleigh temperature lidar for one day (March 17, 2011). These profiles were obtained as a part of the OZITOS II Campaign described above. In each measurement the Rayleigh temperature profiles were compared with the radiosounding aboard the balloon sonde. The effective region for the comparison can be established due figure 2 in the region between 10 up to 30 km aprox. Both instruments have shown good agreement in this region, with a typical relative error lower than 4 %. We have observed also that in this night the three lidar profiles are similar, indicating that the atmospheric conditions were stable. As a comparison with other instrument as the NCEP data and MLS instrument aboard the AURA-NASA satellite has been to do it. It measurements were superposed in the profiles b) and c) (figure 2) showing very good agreement in the region above 20 km. For the region below both measurements (NCEP data and MLS) indicate a discrepancy very similar when are compared with the radiosounding and temperature lidar profiles.

ACKNOWLEDGMENTS

The authors would like to thank JICA (Japan International Cooperation Agency) by financial support of UVO3 Patagonia Project; the CNRS in France for their collaboration in facilitating the shelter and part of the electronic instruments of DIAL.

Analyses used in this paper were produced with the Giovanni online data system, developed and maintained by the NASA GES DISC.

REFERENCES


